

BAROCLINIC AND DIABATIC REGULATION OF THE 10-12 NOVEMBER 2001 SUPERSTORM IN THE BALEARICS

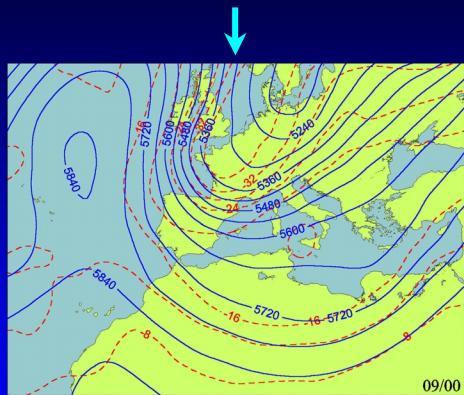
European Conference on Severe Storms 2002 (Prague, Czech Republic, 26-30 August)

R. Romero, V. Homar, C. Ramis, S. Alonso

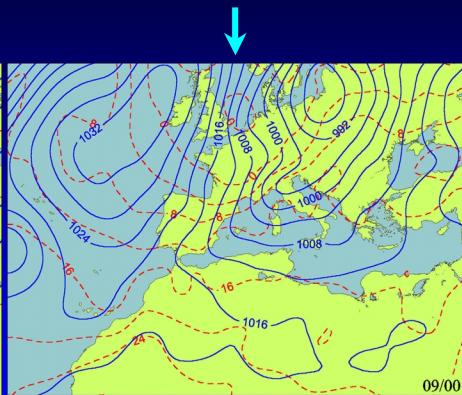


NCEP-BASED METEOROLOGICAL ANALYSES

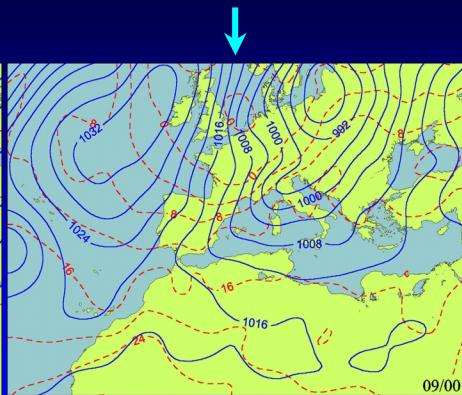
Mid-Upper levels (H 500 / T 500)

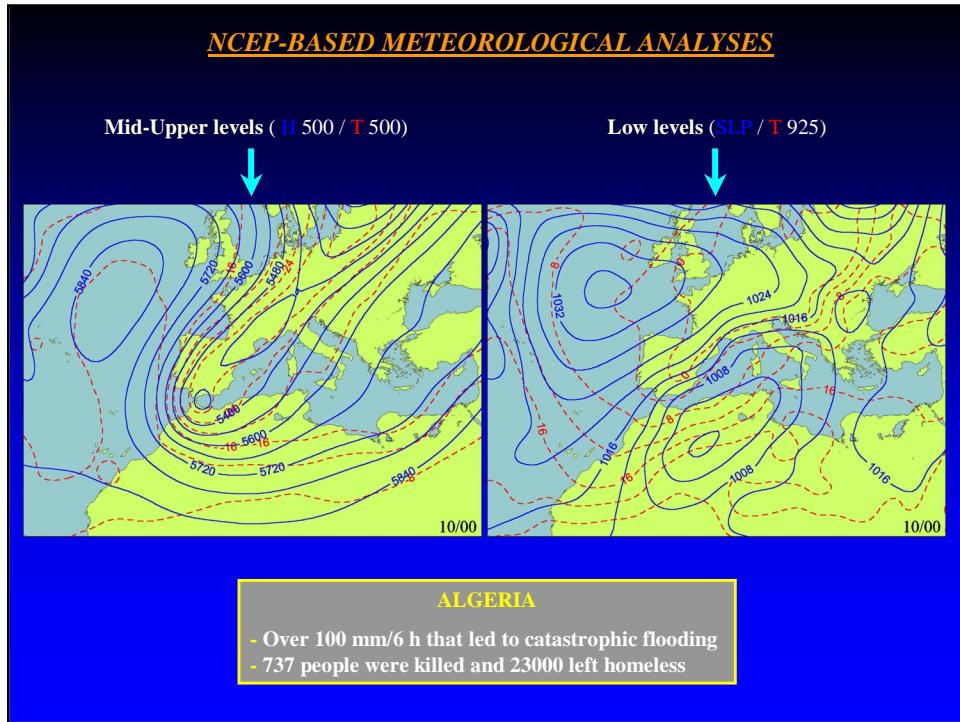
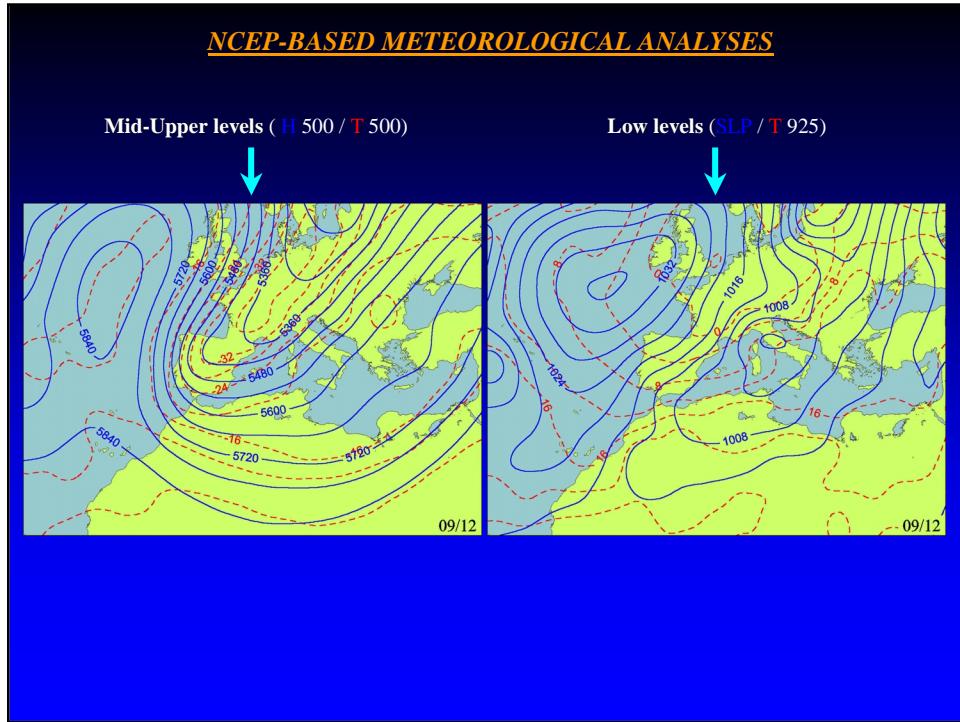


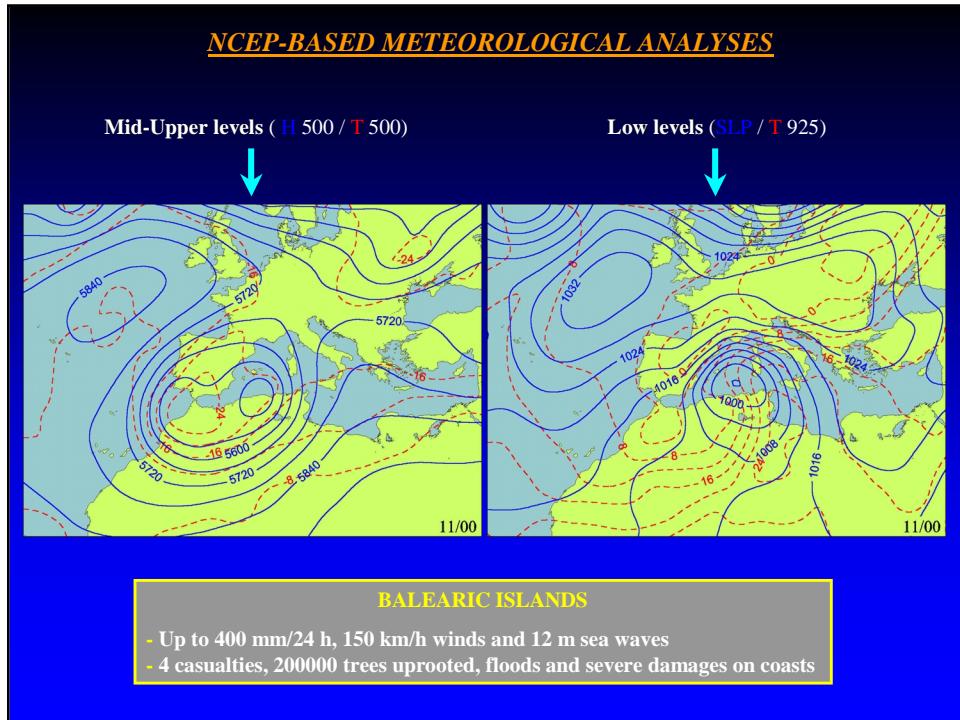
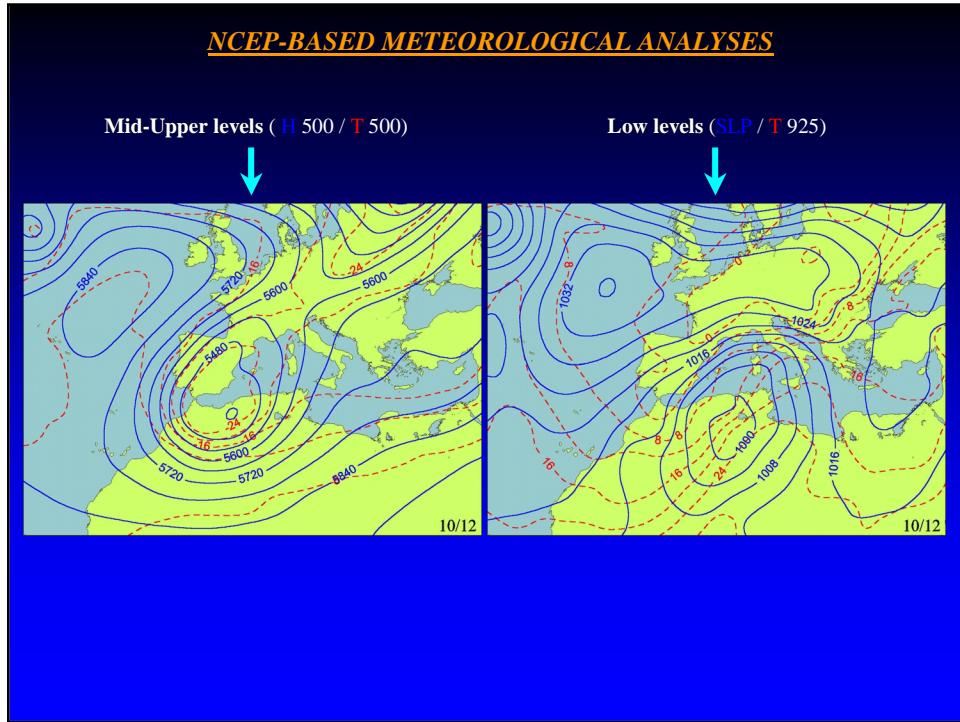
Low levels (SLP / T 925)

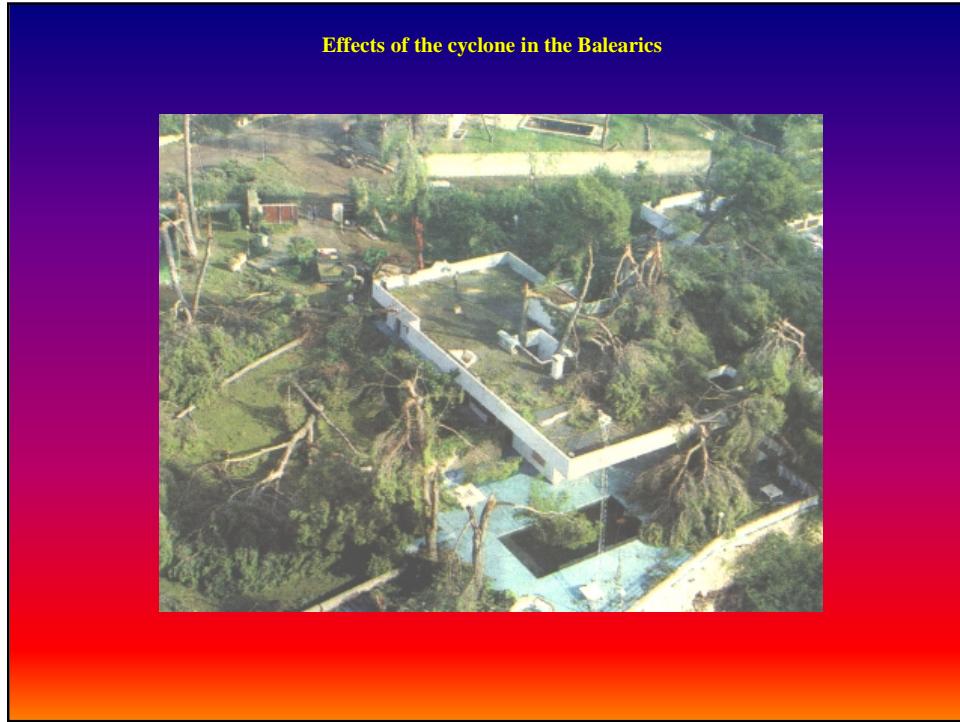
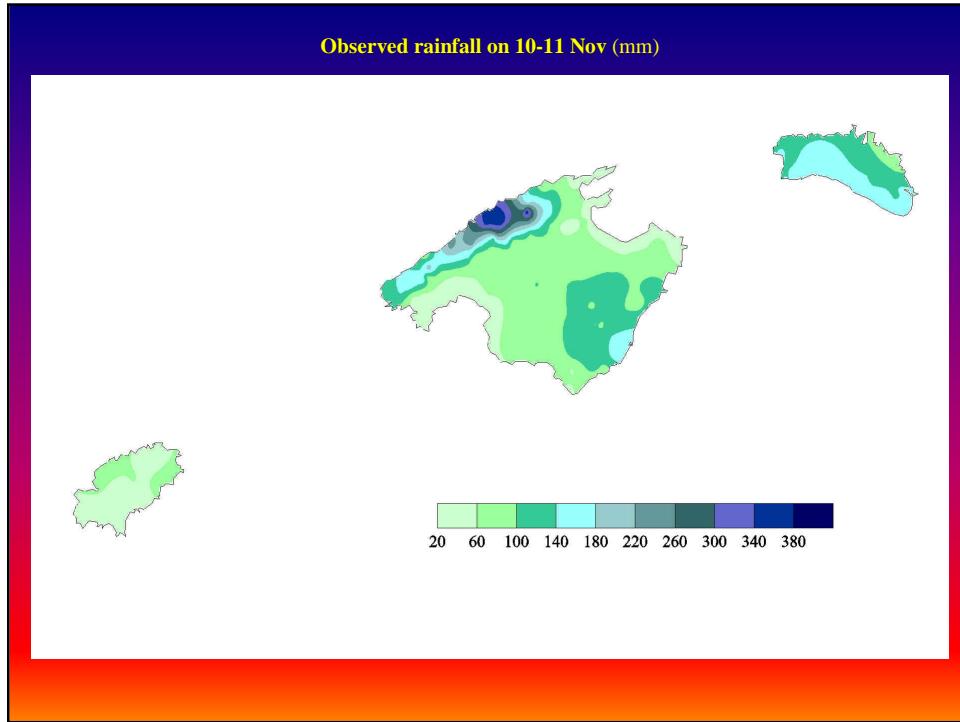


Low levels (SLP / T 925)







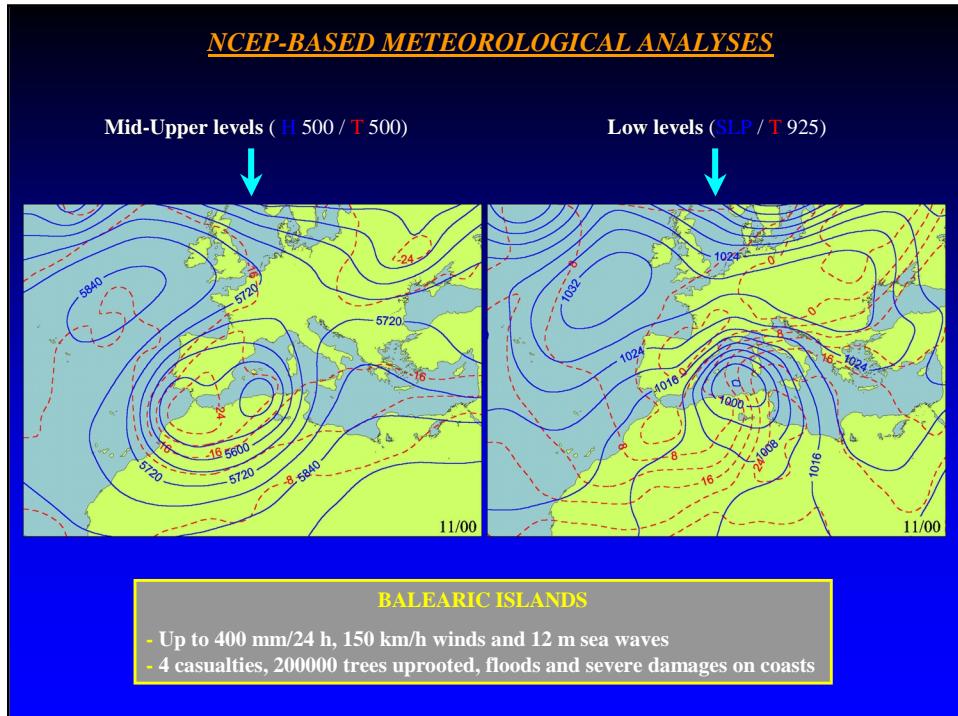
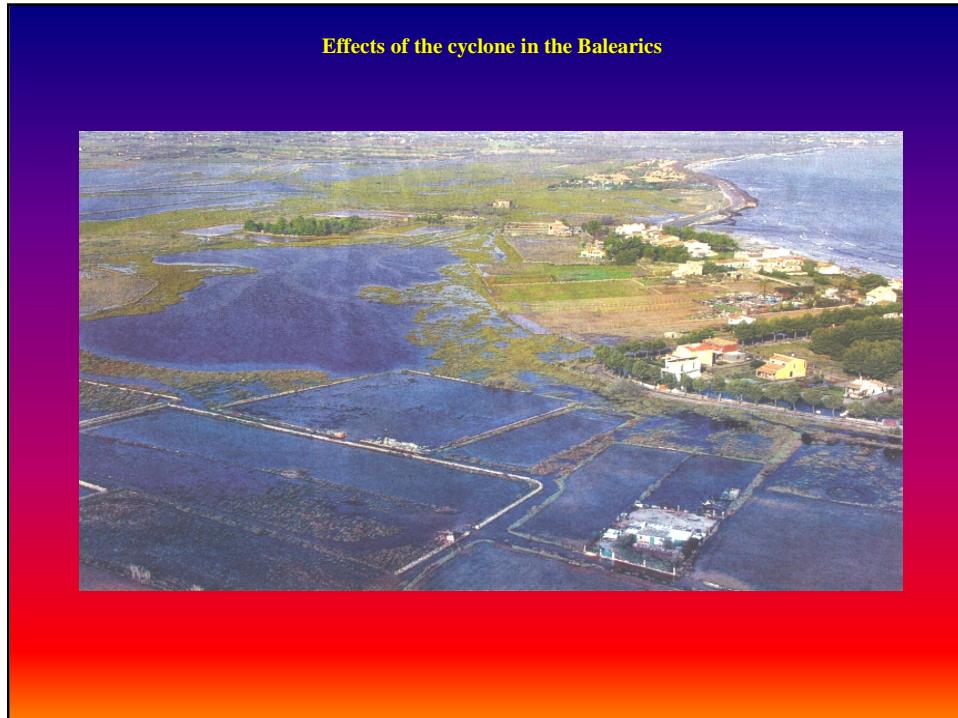


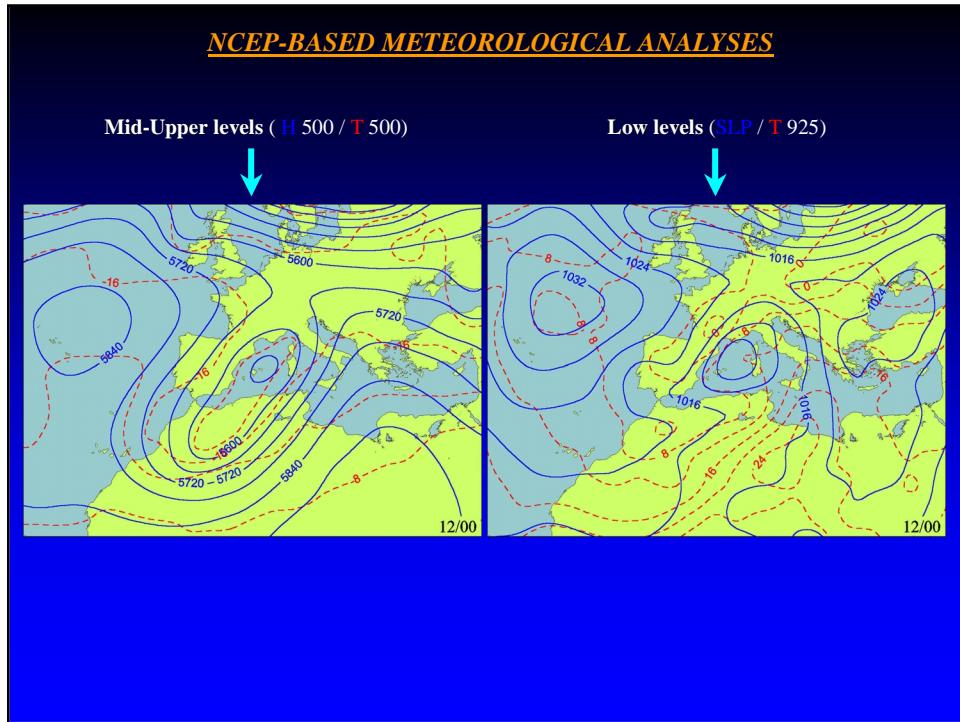
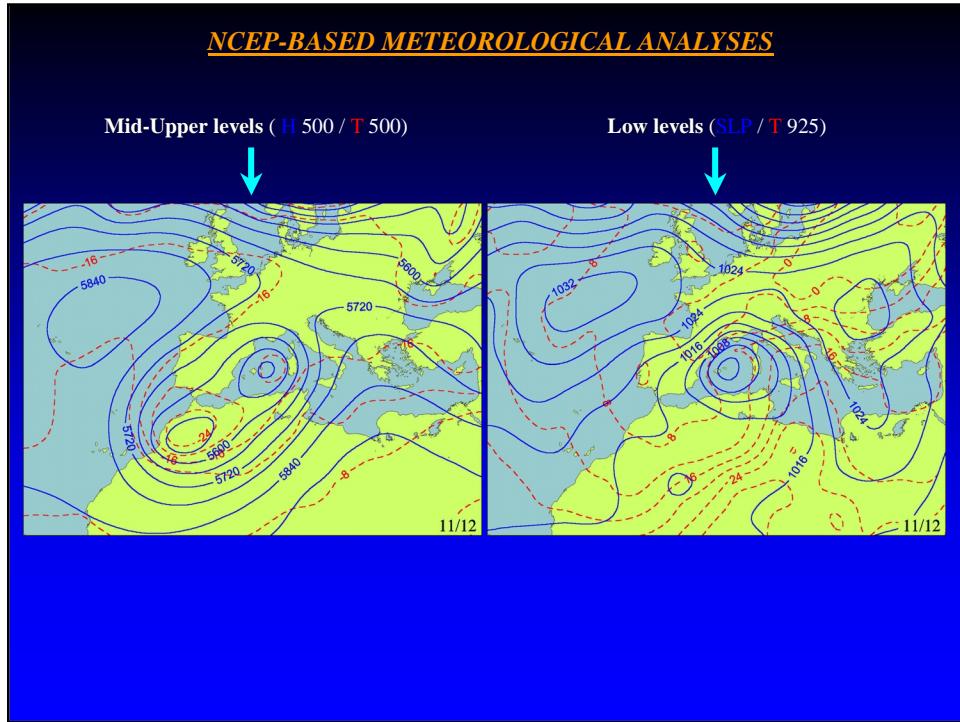
Effects of the cyclone in the Balearics



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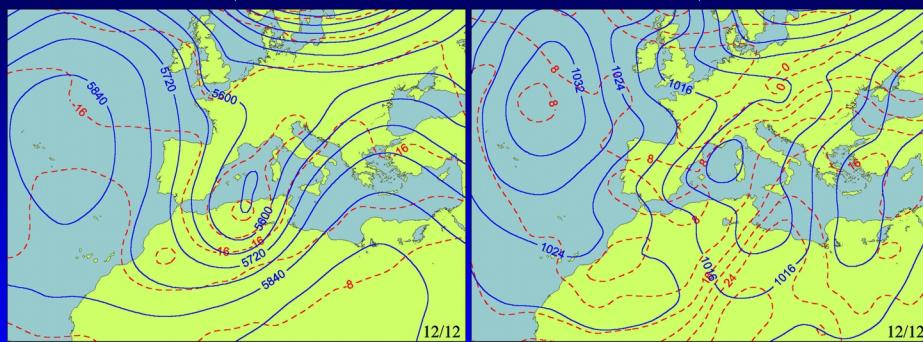




NCEP-BASED METEOROLOGICAL ANALYSES

Mid-Upper levels (H 500 / T 500)

Low levels (SLP / T 925)

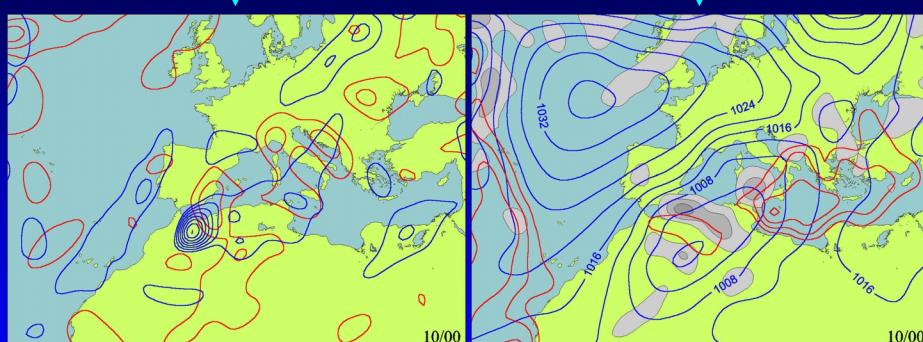


Strong baroclinic development

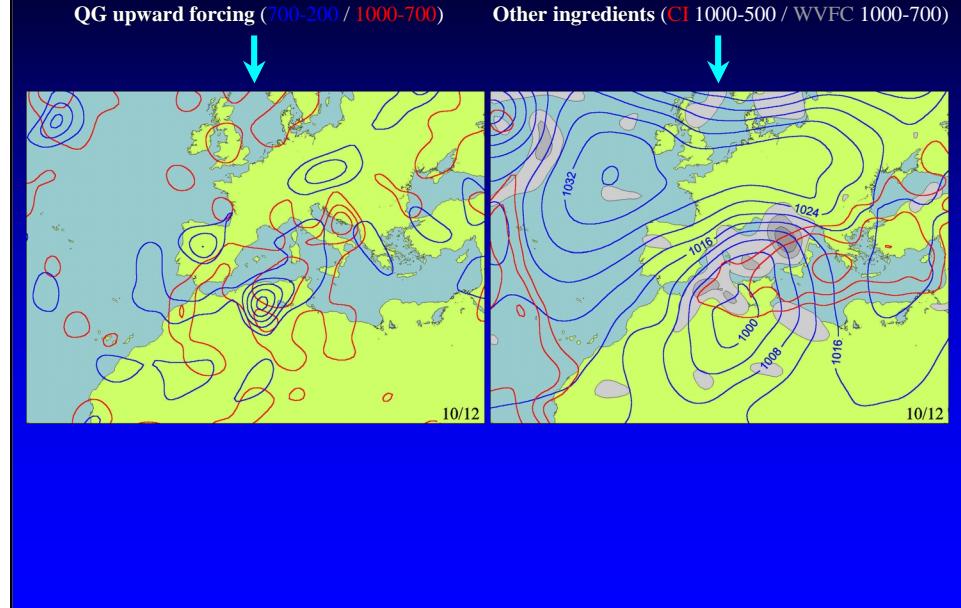
DIAGNOSIS OF THE EVENT (TRADITIONAL)

QG upward forcing (700-200 / 1000-700)

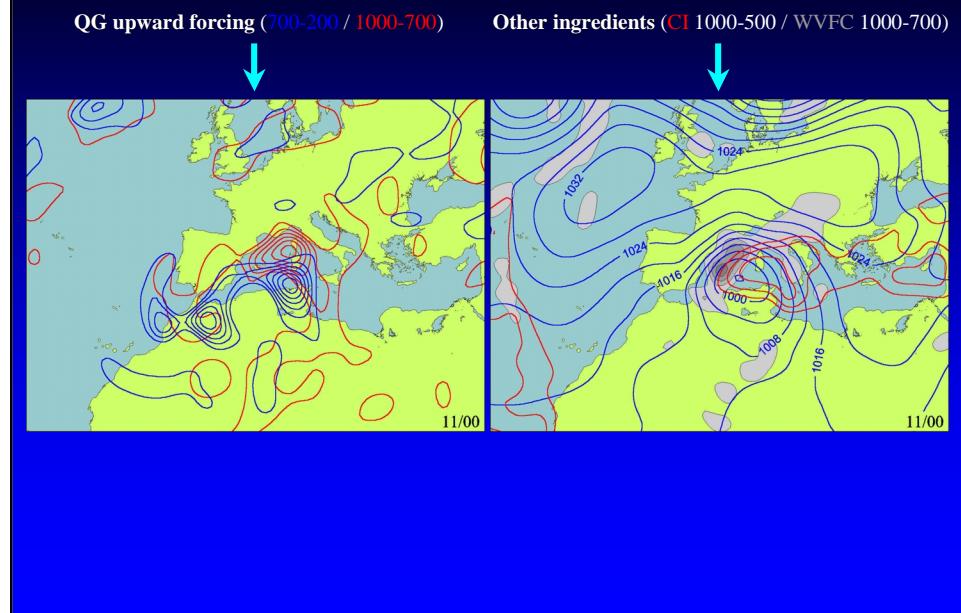
Other ingredients (C1 1000-500 / WVFC 1000-700)



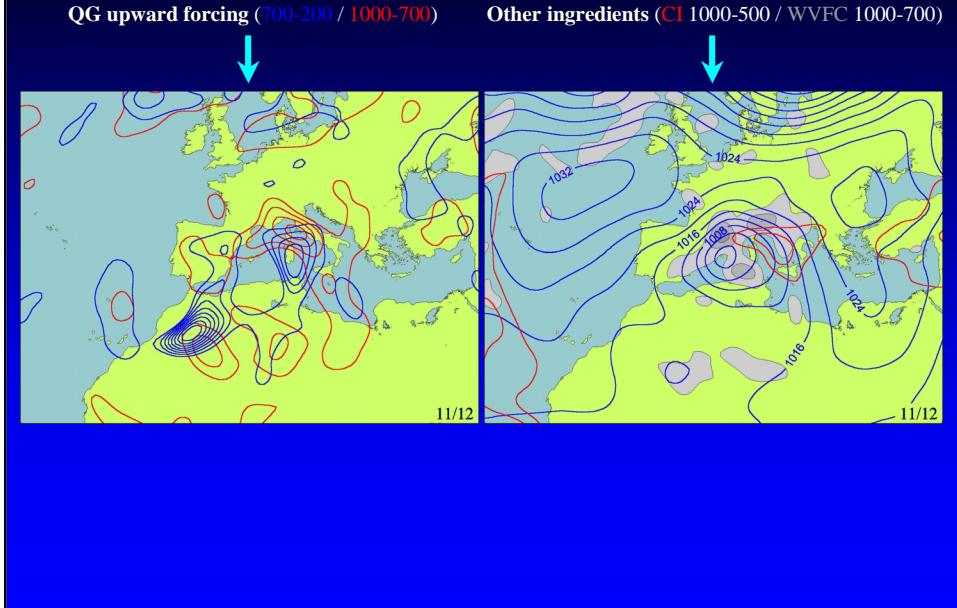
DIAGNOSIS OF THE EVENT (TRADITIONAL)



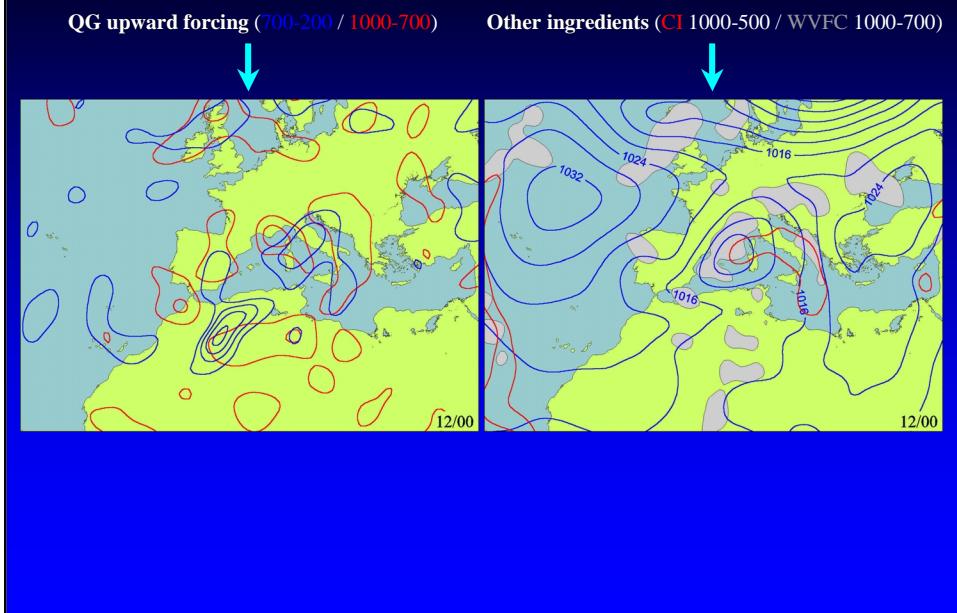
DIAGNOSIS OF THE EVENT (TRADITIONAL)

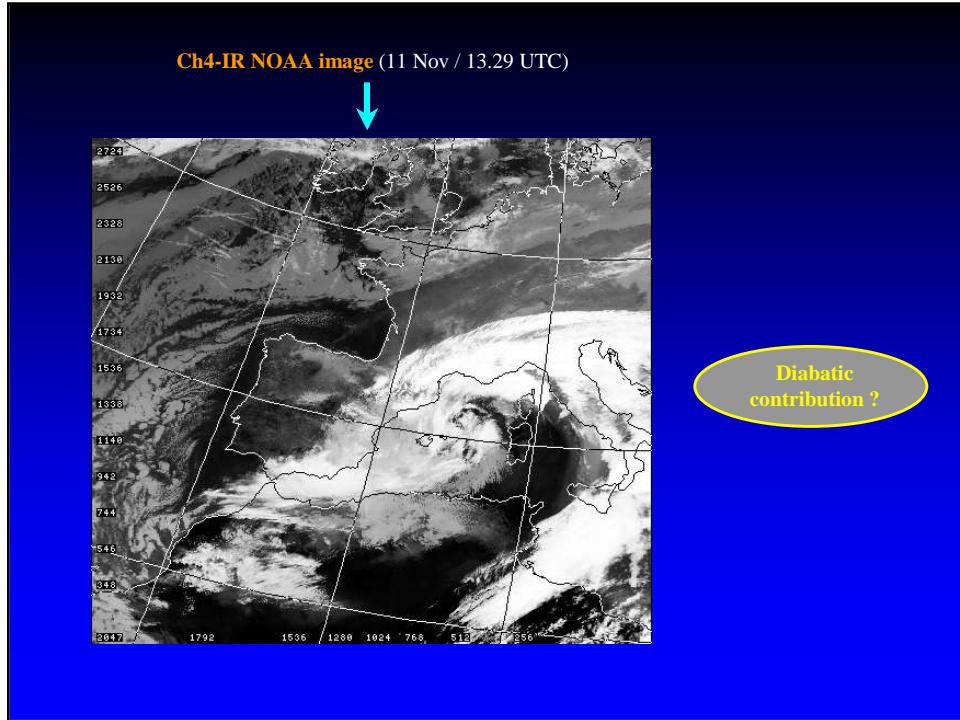
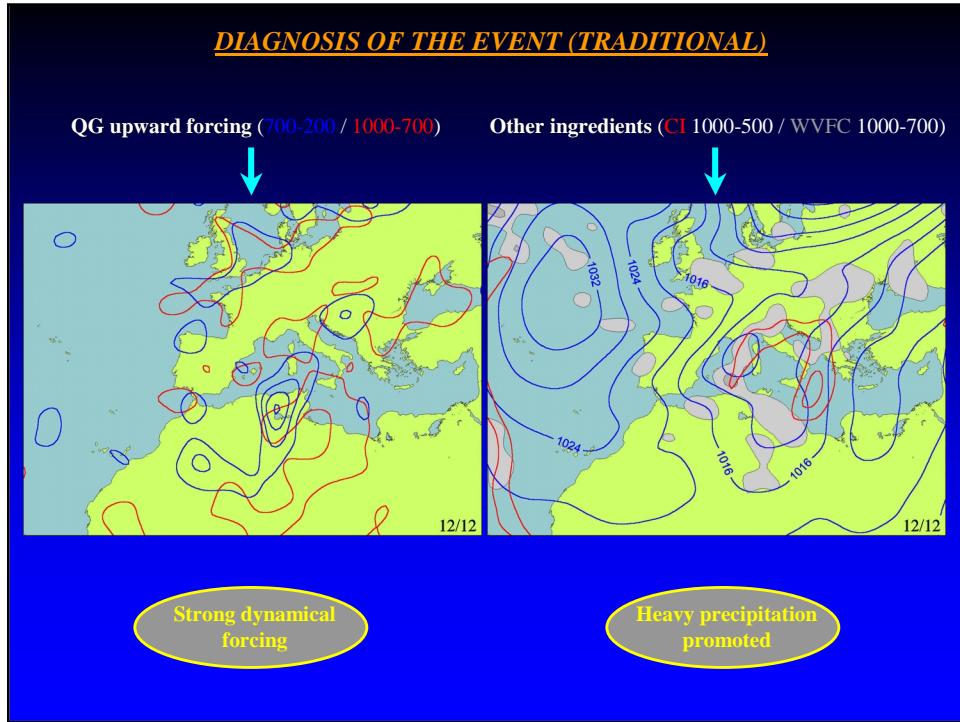


DIAGNOSIS OF THE EVENT (TRADITIONAL)



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PIECEWISE PV INVERSION TECHNIQUE

(Davis and Emanuel; MWR 1991)

1) Balanced flow (ϕ , ψ) given instantaneous distribution of Ertel's PV (q):

* Charney (1955) nonlinear balance equation

$$\nabla^2\phi = \nabla \cdot f \nabla \psi + 2m^2 \left[\frac{\partial^2\psi}{\partial x^2} \frac{\partial^2\psi}{\partial y^2} - \left(\frac{\partial^2\psi}{\partial x \partial y} \right)^2 \right]$$

f Coriolis parameter *m* map-scale factor

* Approximate form of Ertel's PV

$$q = \frac{g\kappa\pi}{p} \left[(f + m^2 \nabla^2\psi) \frac{\partial^2\phi}{\partial \pi^2} - m^2 \left(\frac{\partial^2\psi}{\partial x \partial \pi} \frac{\partial^2\phi}{\partial x \partial \pi} + \frac{\partial^2\psi}{\partial y \partial \pi} \frac{\partial^2\phi}{\partial y \partial \pi} \right) \right]$$

p pressure *g* gravity $\kappa = R_d/C_p$ $\pi = C_p(p/p_0)^\kappa$

* B.C Lateral (Dirichlet) / Top and Bottom (Neumann):

$$\frac{\partial\phi}{\partial\pi} = f \frac{\partial\psi}{\partial\pi} = -\theta$$

θ potential temperature

2) Reference state: Balanced flow ($\bar{\phi}$, $\bar{\psi}$) given time mean distribution of Ertel's PV (\bar{q}):

* Same equations as in 1), except using time mean fields instead of instantaneous fields

3) Perturbation fields (ϕ' , ψ' , q') defined through: $(q, \phi, \psi) = (\bar{q}, \bar{\phi}, \bar{\psi}) + (q', \phi', \psi')$

PIECEWISE PV INVERSION TECHNIQUE

4) We consider that q' is partitioned into N portions or anomalies: $q' = \sum_{n=1}^N q_n$

5) Piecewise inversion: (ϕ_n, ψ_n) associated with q_n ?

...After substitution of the above summations in the
balance and PV equations and some rearrangements
of the nonlinear terms:

$$\begin{aligned} \phi' &= \sum_{n=1}^N \phi_n \\ \psi' &= \sum_{n=1}^N \psi_n \end{aligned}$$

$$\nabla^2\phi_n = \nabla \cdot f \nabla \psi_n + 2m^2 \left(\frac{\partial^2\psi^*}{\partial x^2} \frac{\partial^2\psi_n}{\partial y^2} + \frac{\partial^2\psi^*}{\partial y^2} \frac{\partial^2\psi_n}{\partial x^2} - 2 \frac{\partial^2\psi^*}{\partial x \partial y} \frac{\partial^2\psi_n}{\partial y \partial x} \right)$$

$$\begin{aligned} q_n &= \frac{g\kappa\pi}{p} \left[(f + m^2 \nabla^2\psi^*) \frac{\partial^2\phi_n}{\partial \pi^2} + m^2 \frac{\partial^2\phi^*}{\partial \pi^2} \nabla^2\psi_n \right. \\ &\quad \left. - m^2 \left(\frac{\partial^2\phi^*}{\partial x \partial \pi} \frac{\partial^2\psi_n}{\partial x \partial \pi} + \frac{\partial^2\phi^*}{\partial y \partial \pi} \frac{\partial^2\psi_n}{\partial y \partial \pi} \right) - m^2 \left(\frac{\partial^2\psi^*}{\partial x \partial \pi} \frac{\partial^2\phi_n}{\partial x \partial \pi} + \frac{\partial^2\psi^*}{\partial y \partial \pi} \frac{\partial^2\phi_n}{\partial y \partial \pi} \right) \right] \end{aligned}$$

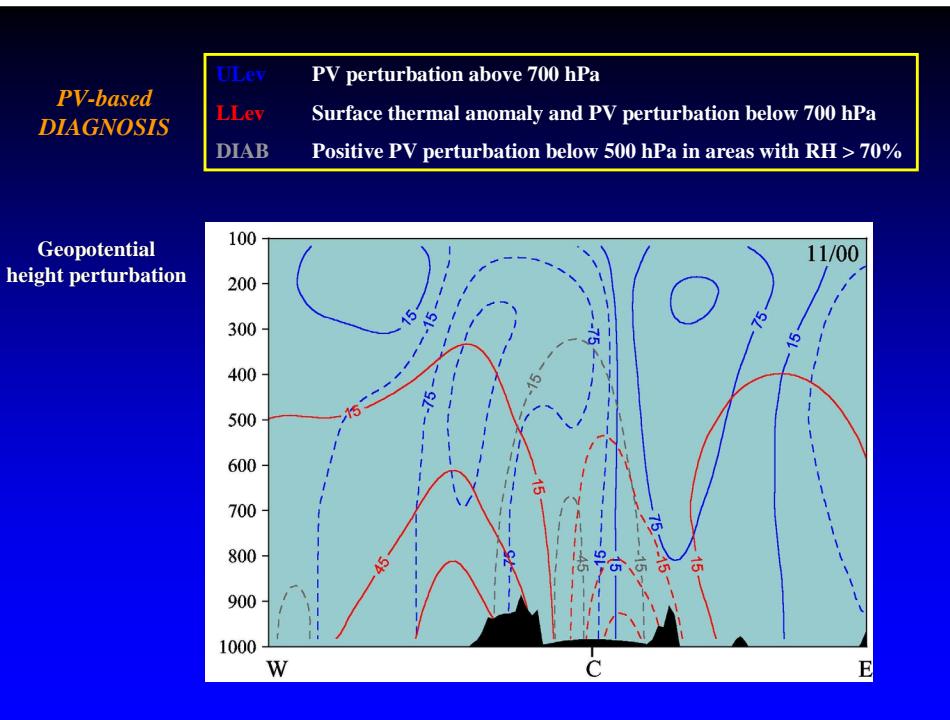
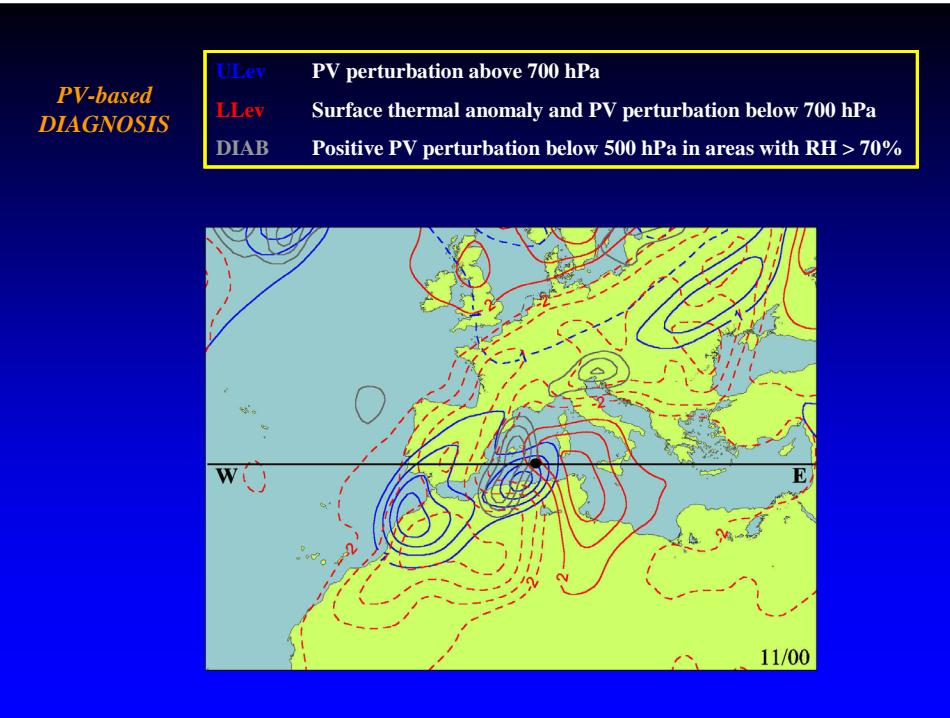
where $(\cdot)^* = \bar{(\cdot)} + \frac{1}{2}(\cdot)'$

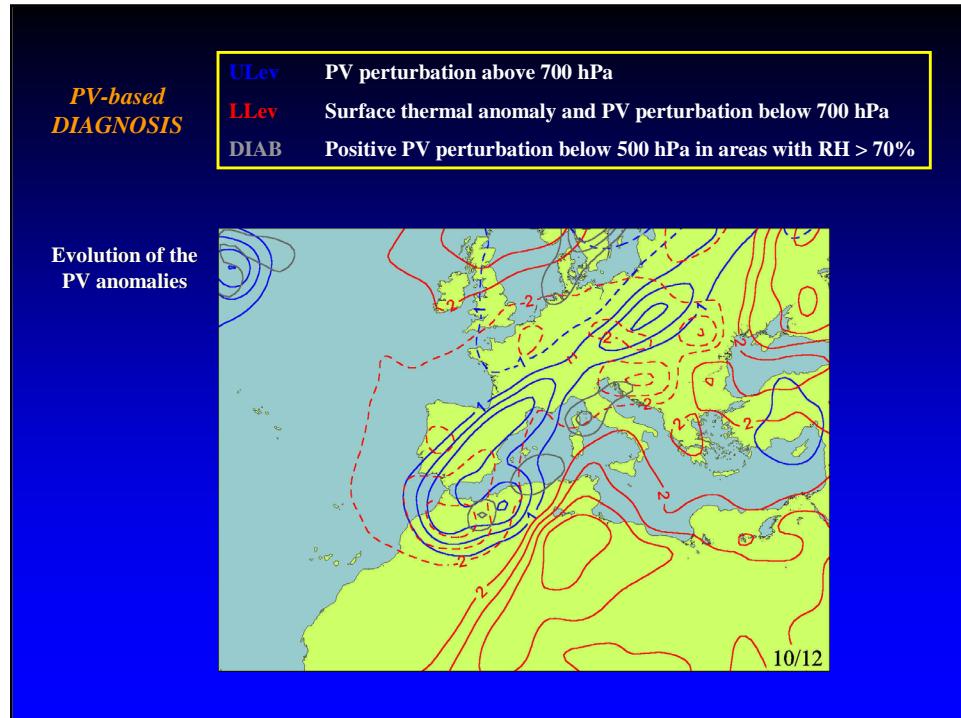
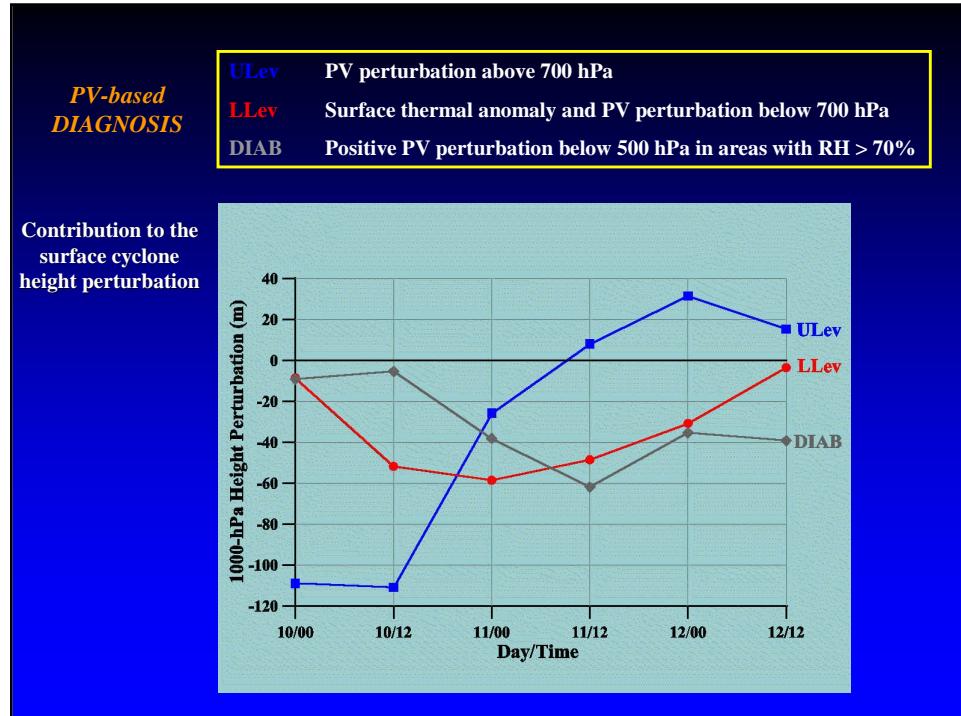
B.C: Lateral (Dirichlet with ϕ_n and ψ_n) / Top and bottom (Neumann with θ_n)

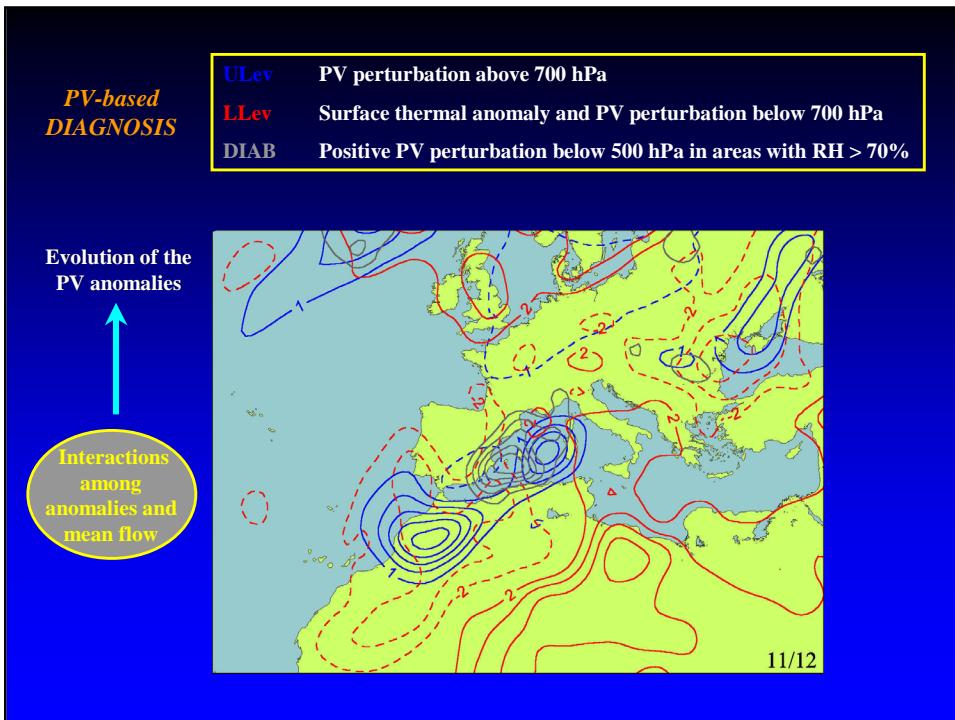
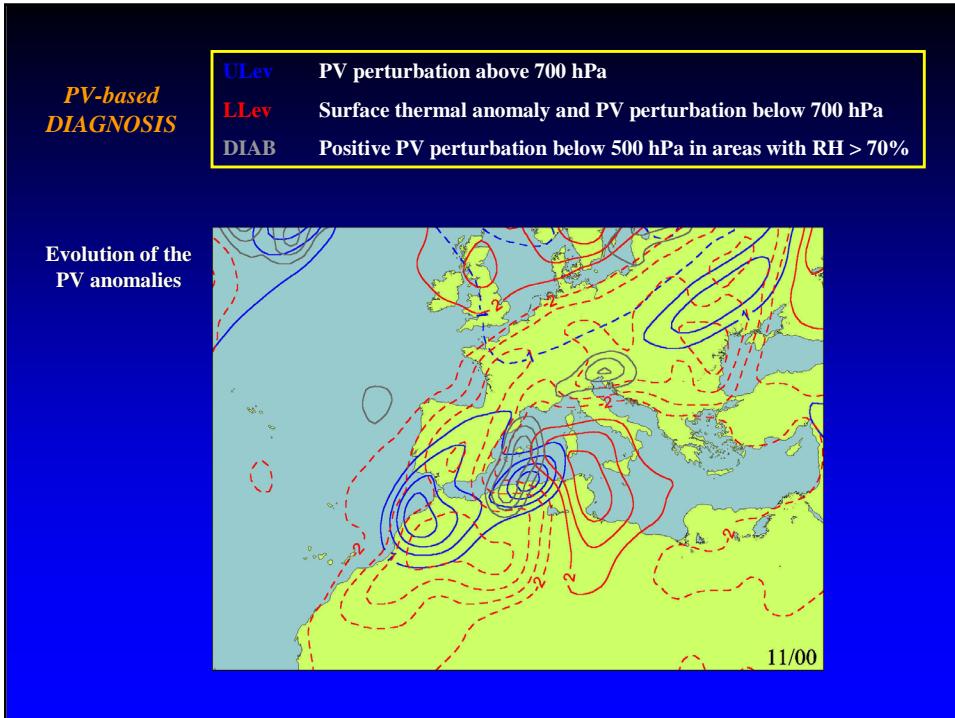
Time interval: 9/00 - 12/12 every 12 h, using the NCEP meteorological analyses

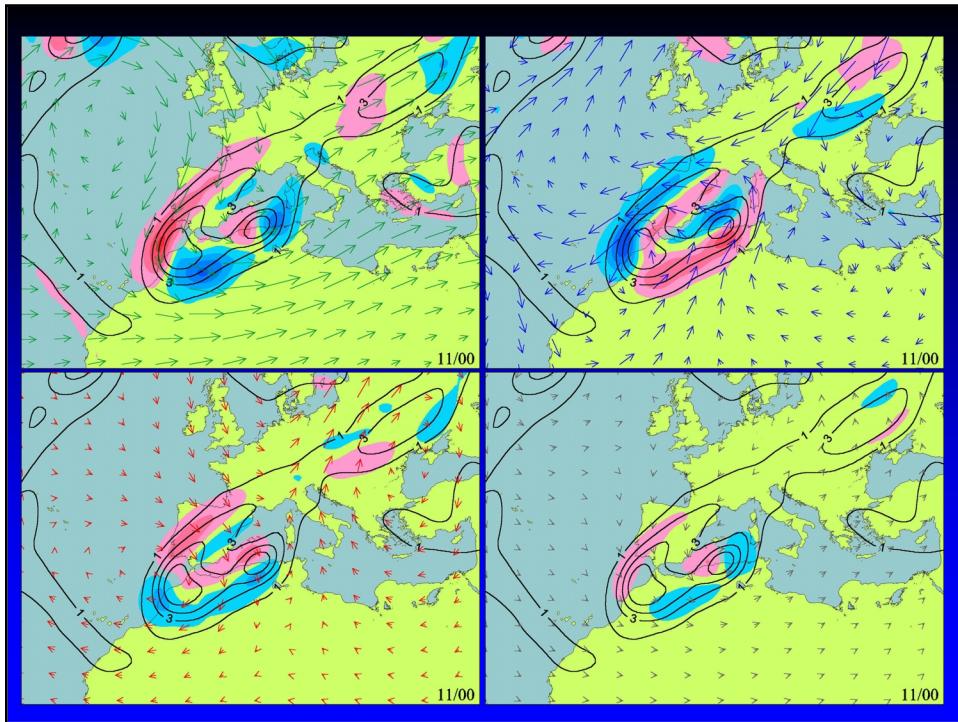
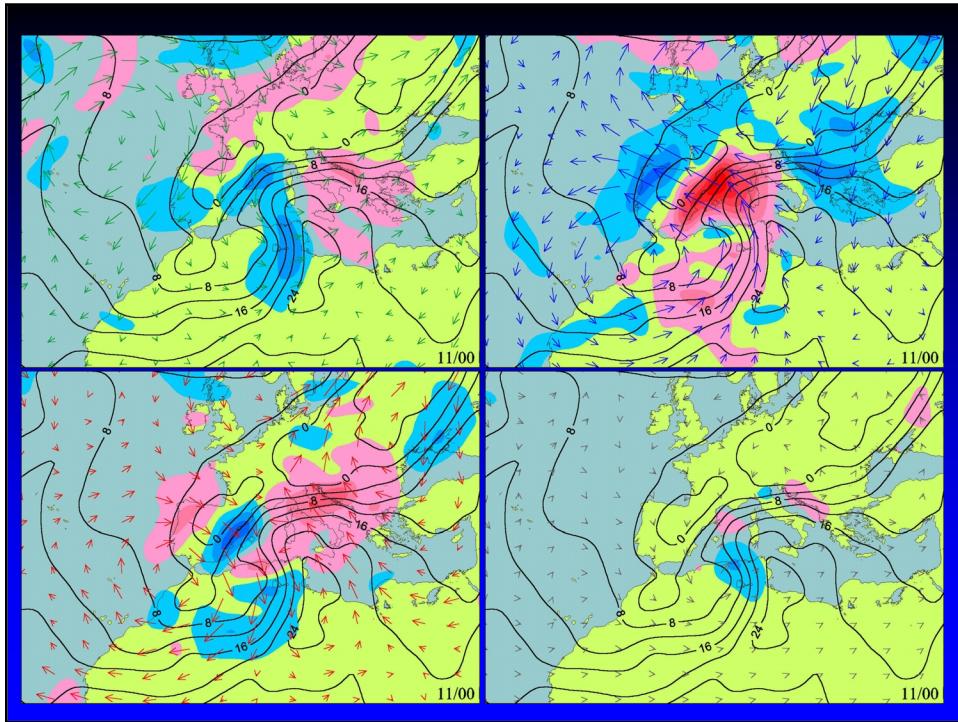
* In our case study: Reference state: 7-day time average for the period 7/00 - 14/00

Anomalies: ULev, LLev, DIAB





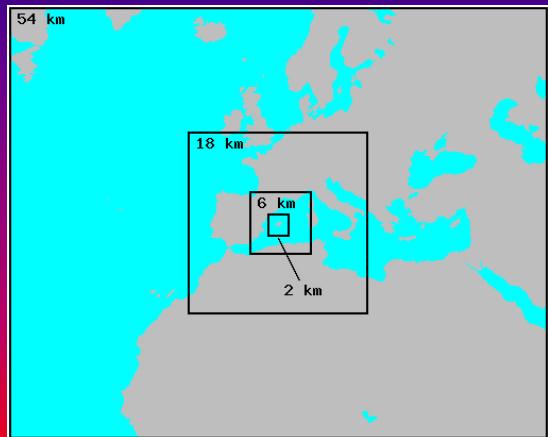




MM5 MODEL SIMULATIONS

- Simulation period (48 h):

10 Nov/00 UTC → 12 Nov/00 UTC



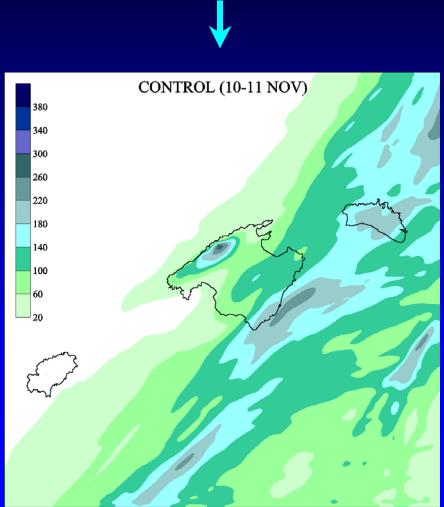
- IC and BC:

2.5° NCEP analyses available at 00/12 UTC
+ upper-air and surface observations

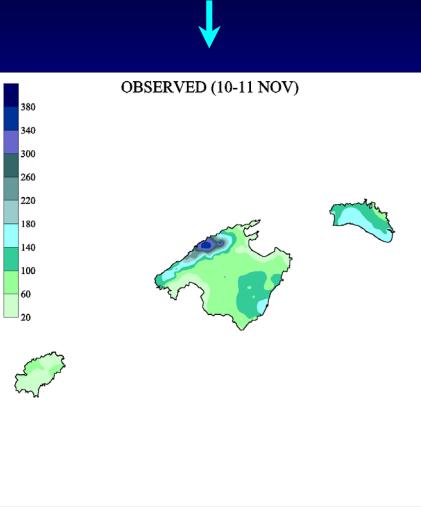
- Domains:

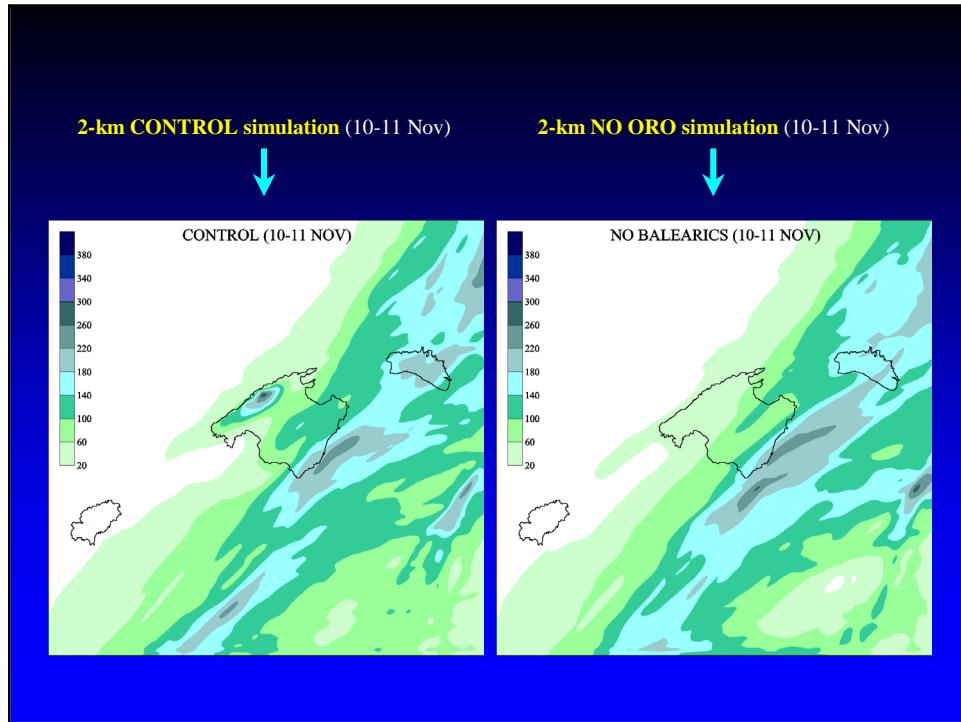
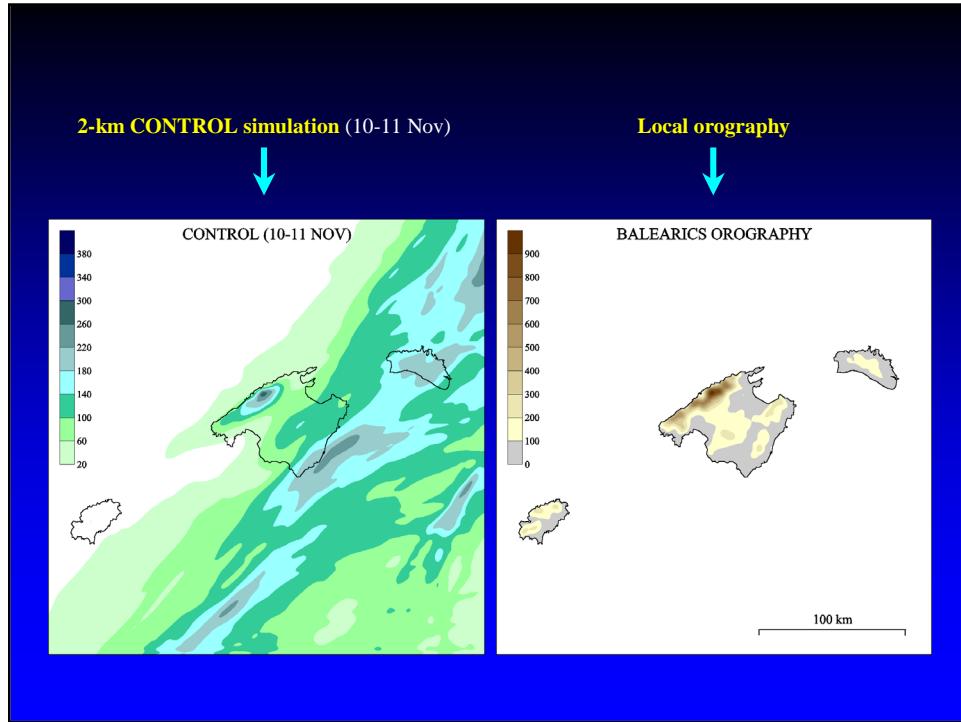
4 domains (54, 18, 6 and 2 km) with 2-way nesting strategy

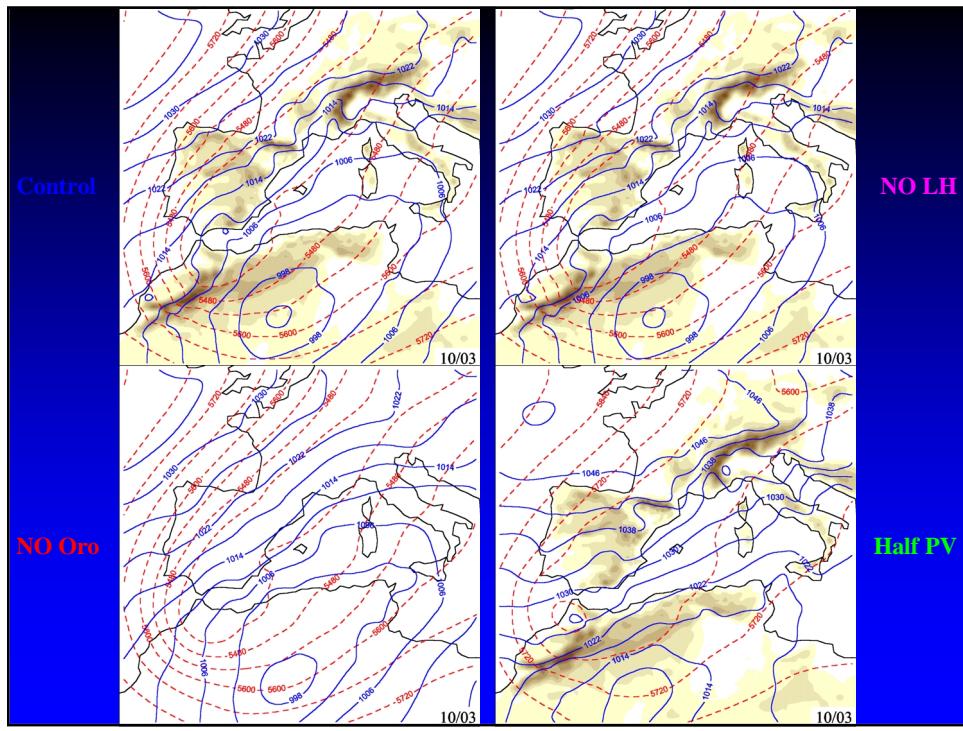
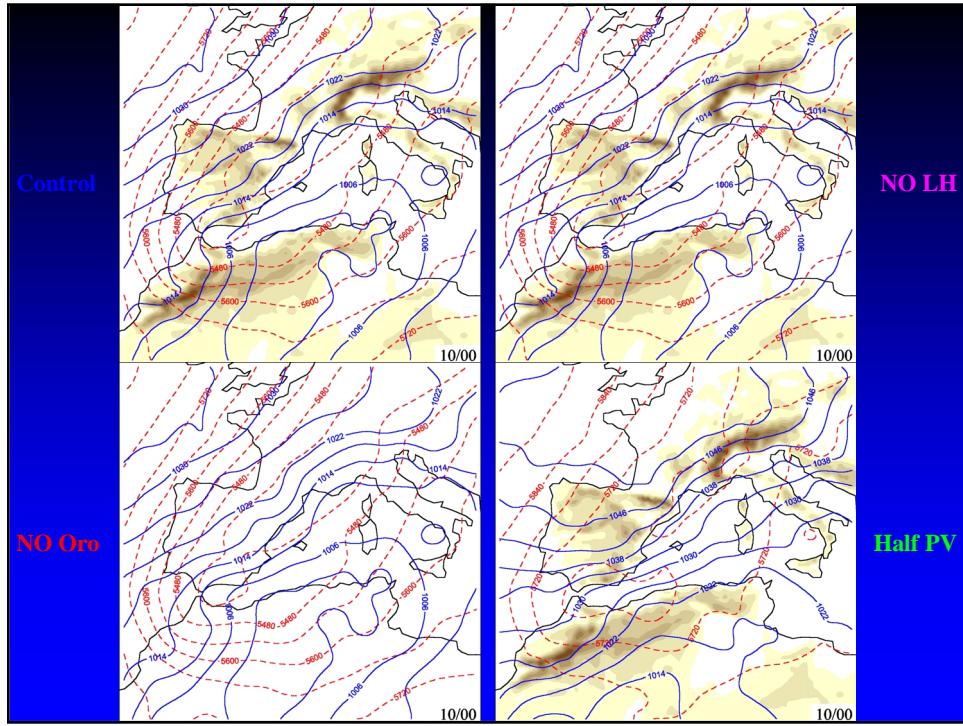
2-km CONTROL simulation (10-11 Nov)

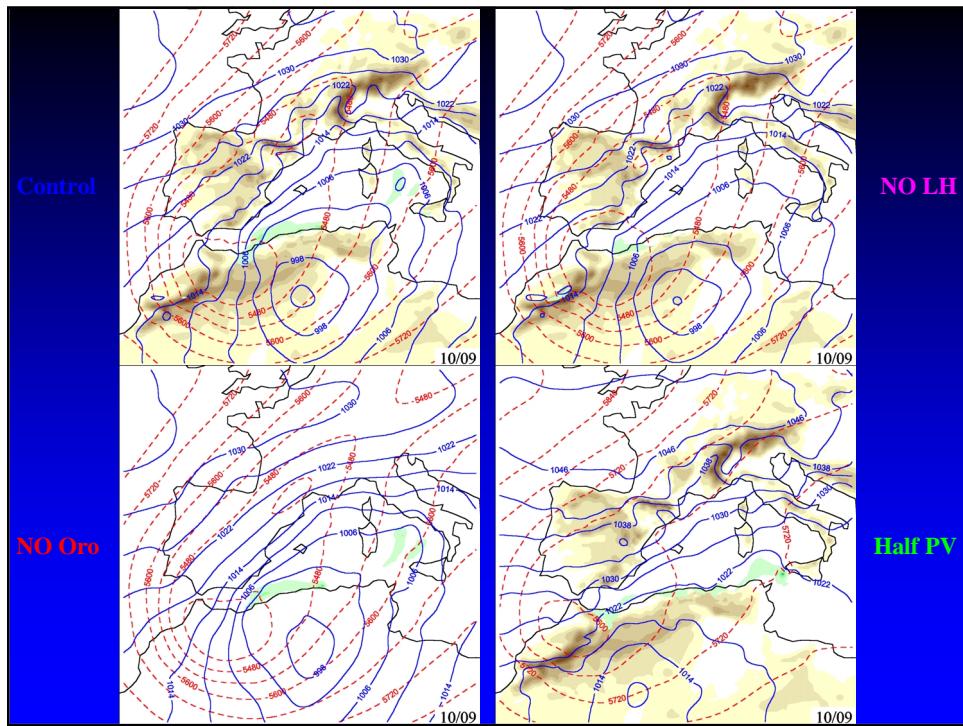
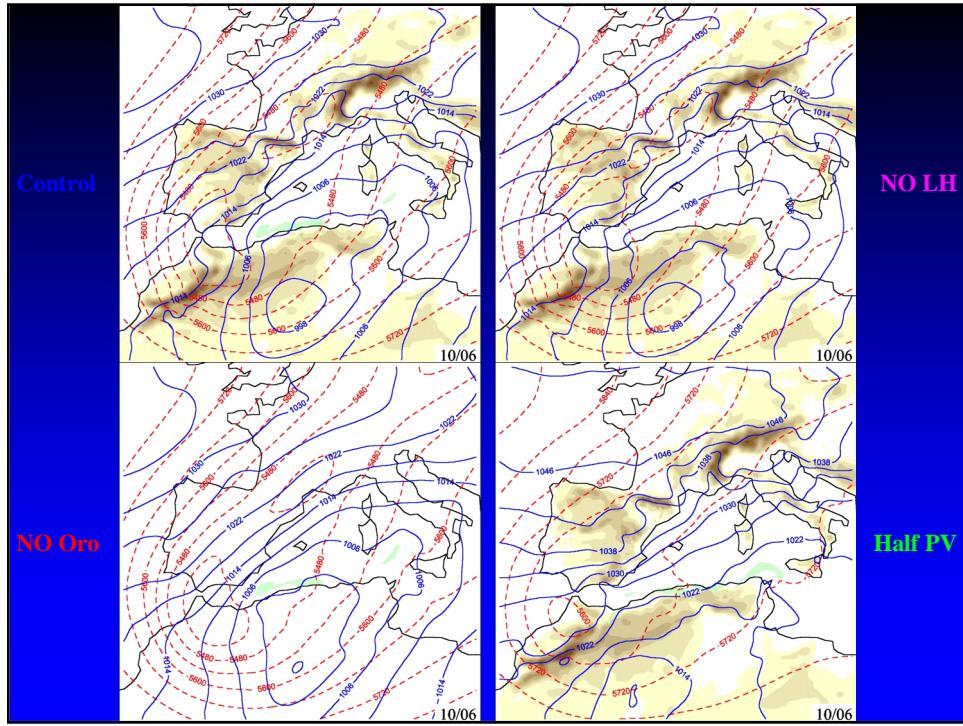


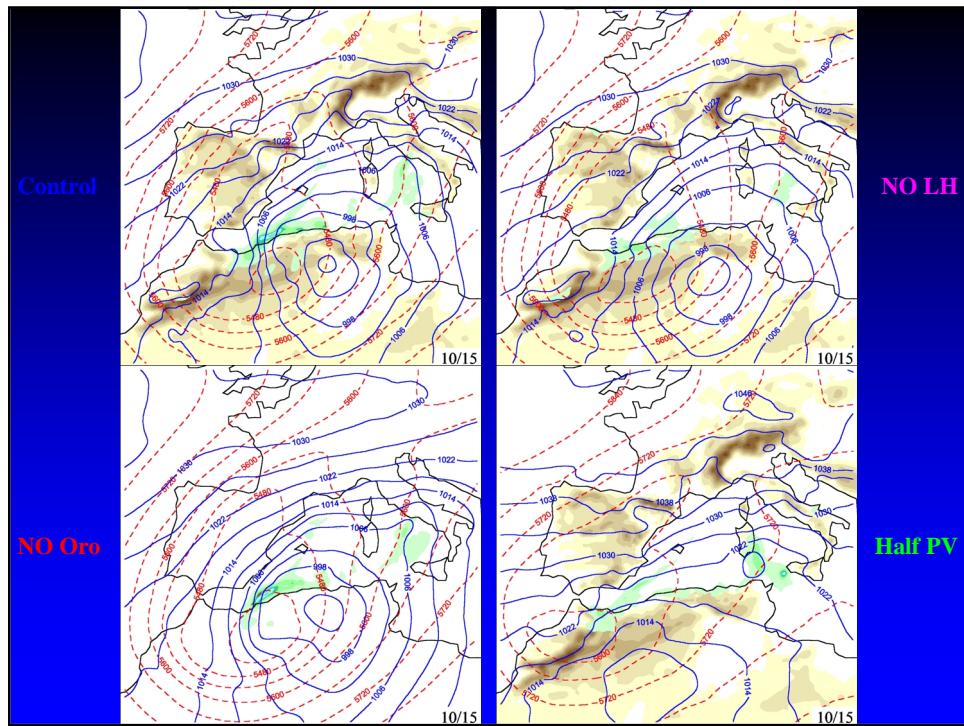
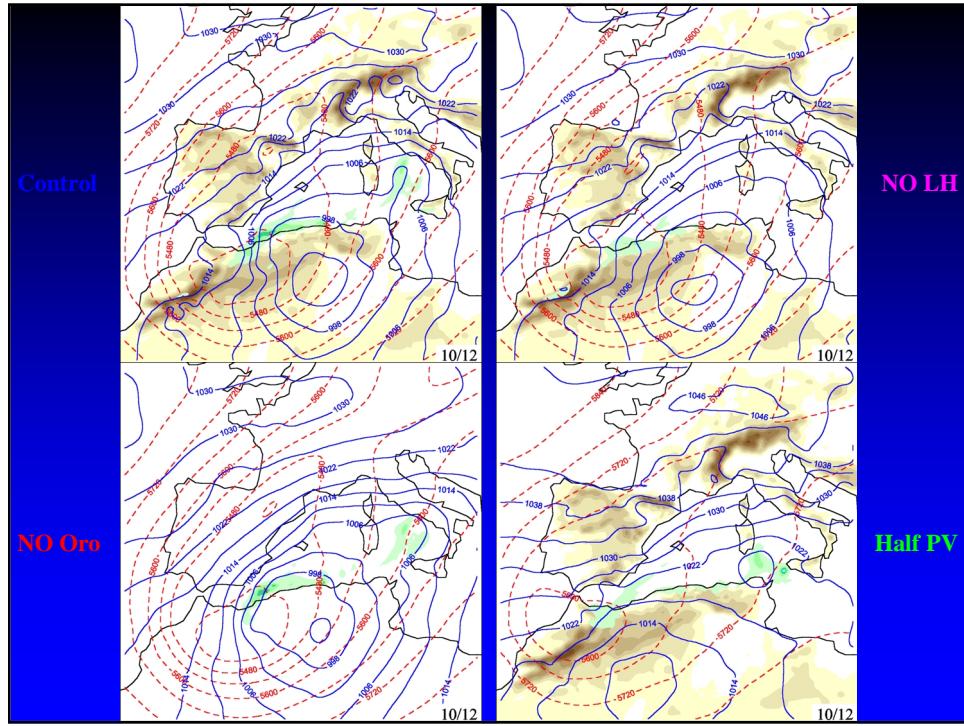
Observed rainfall (10-11 Nov)

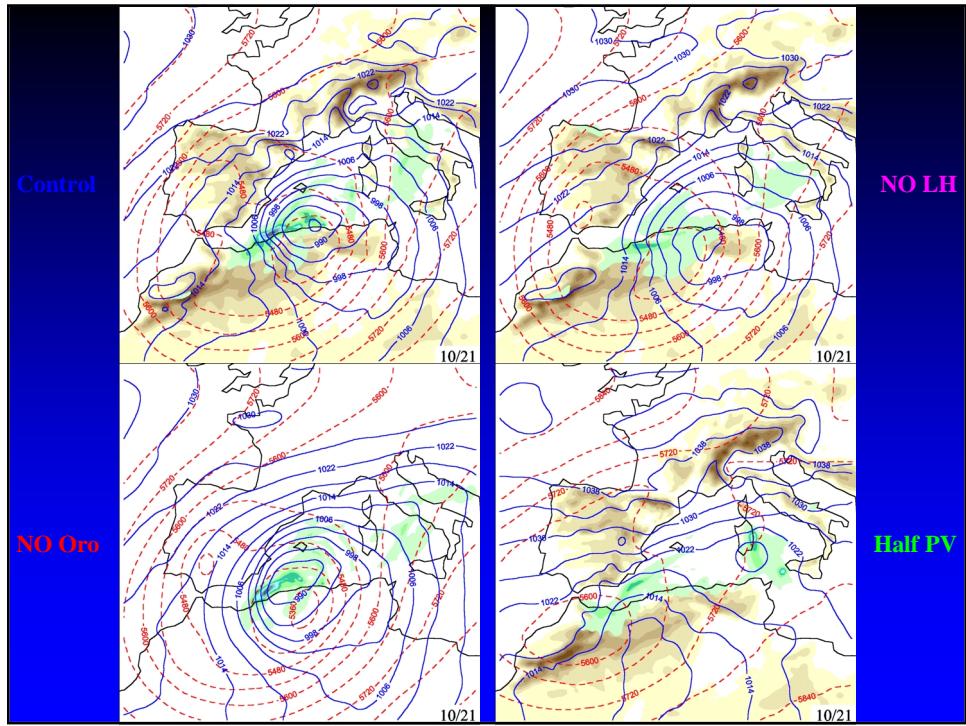
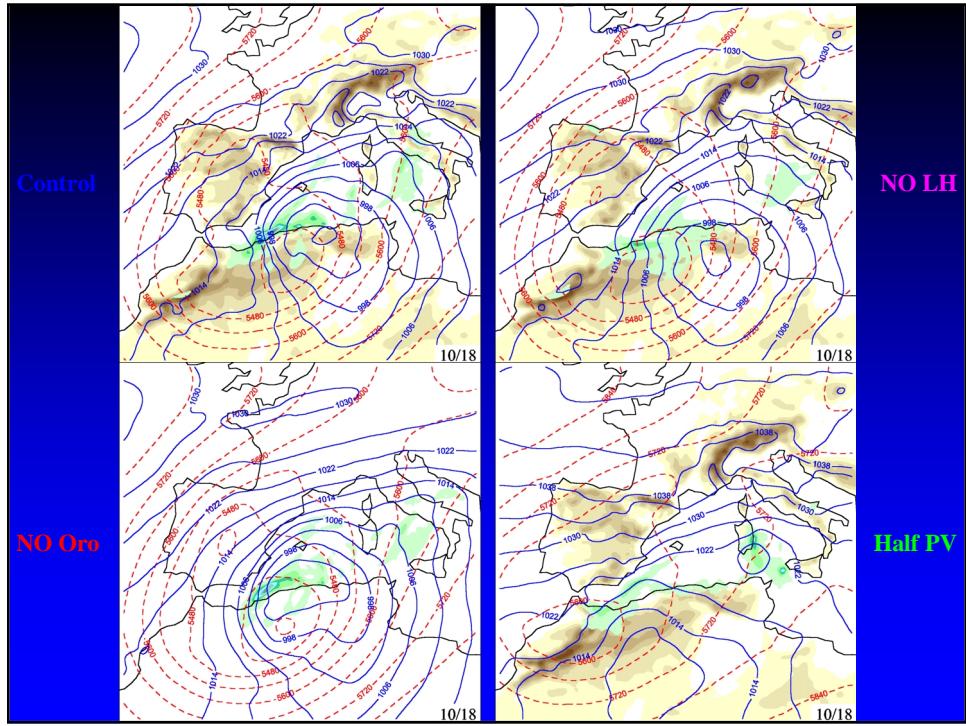


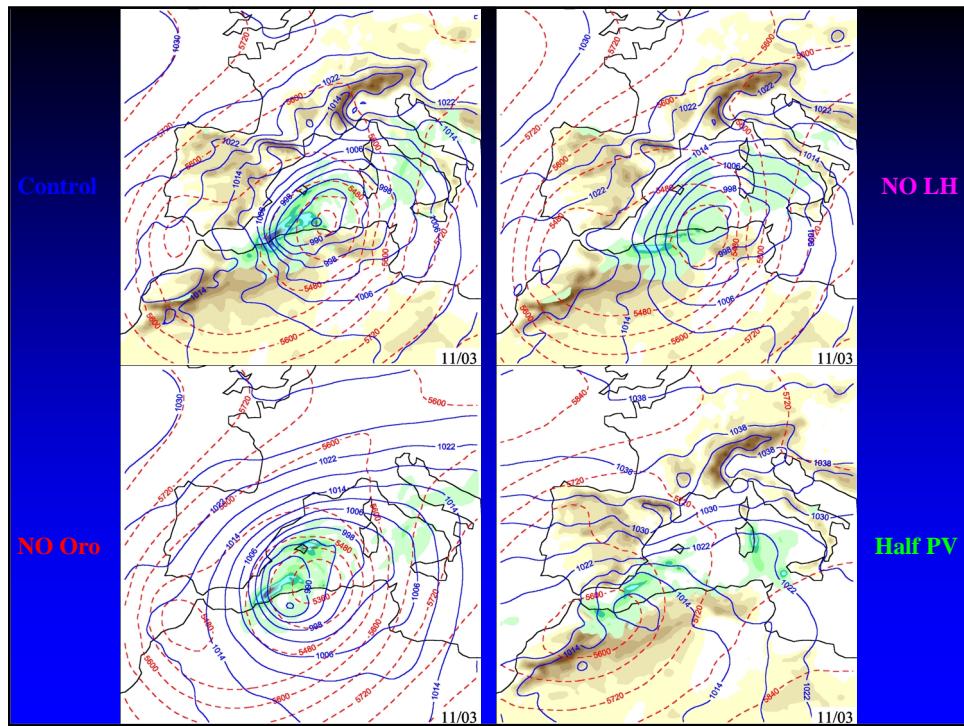
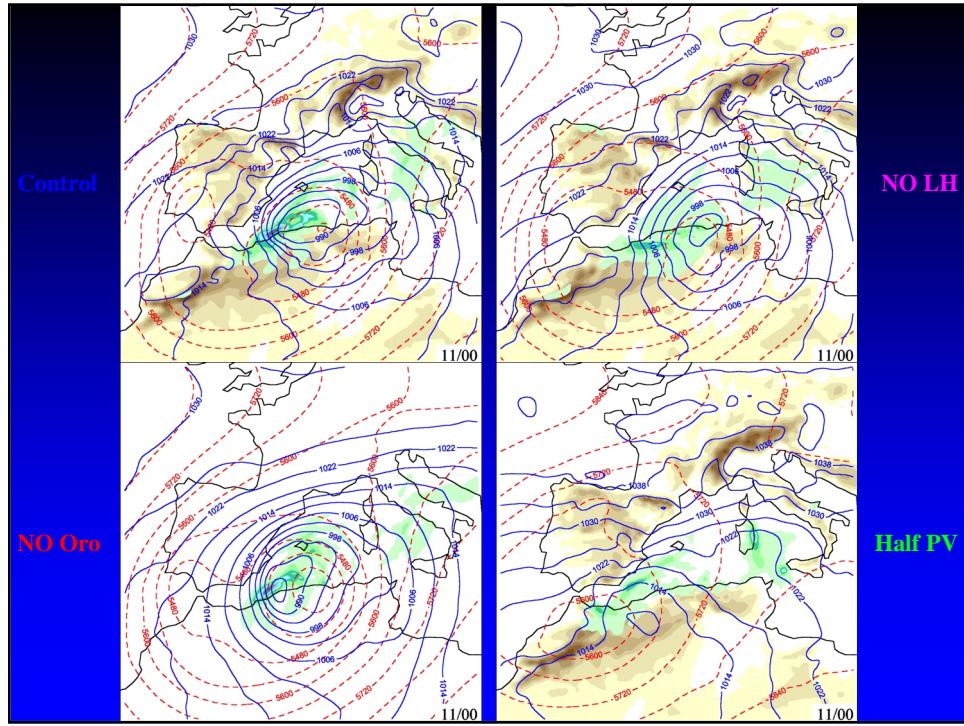


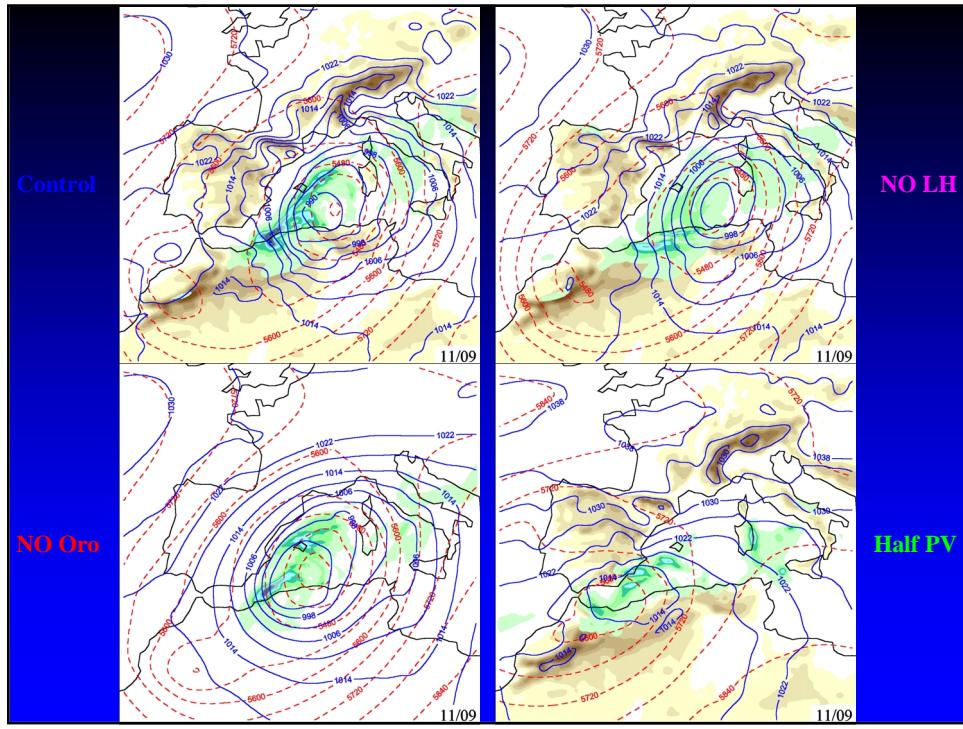
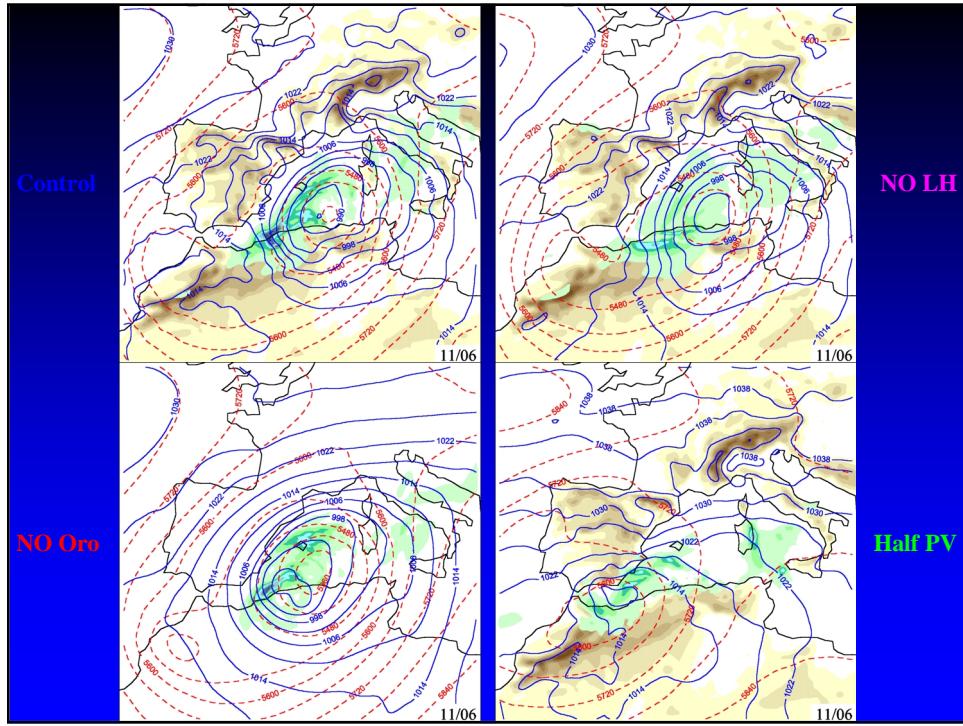


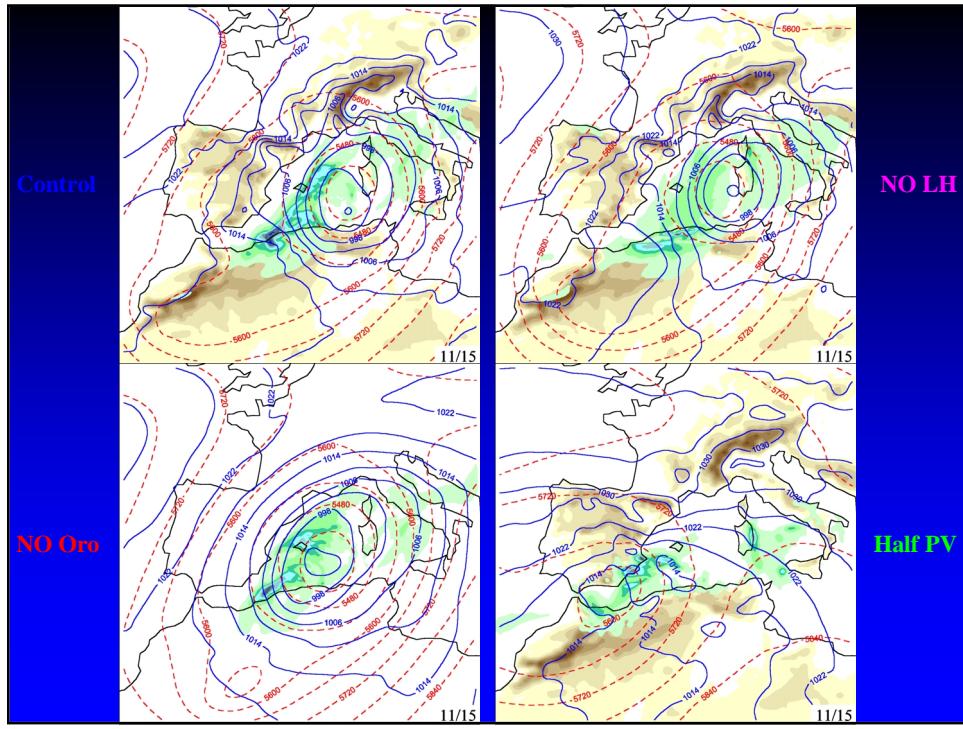
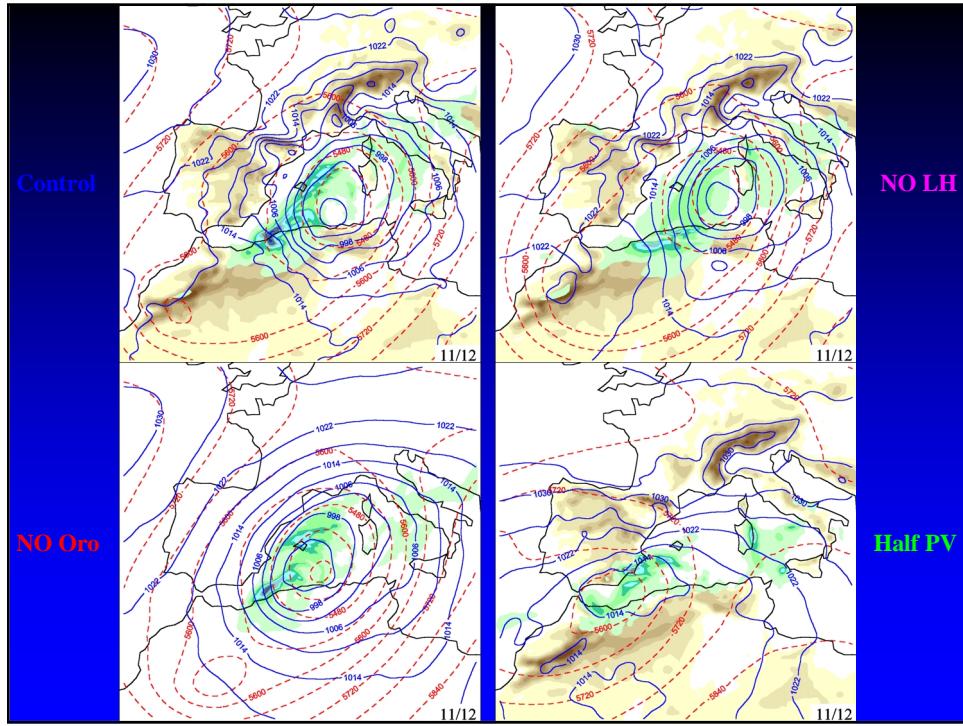


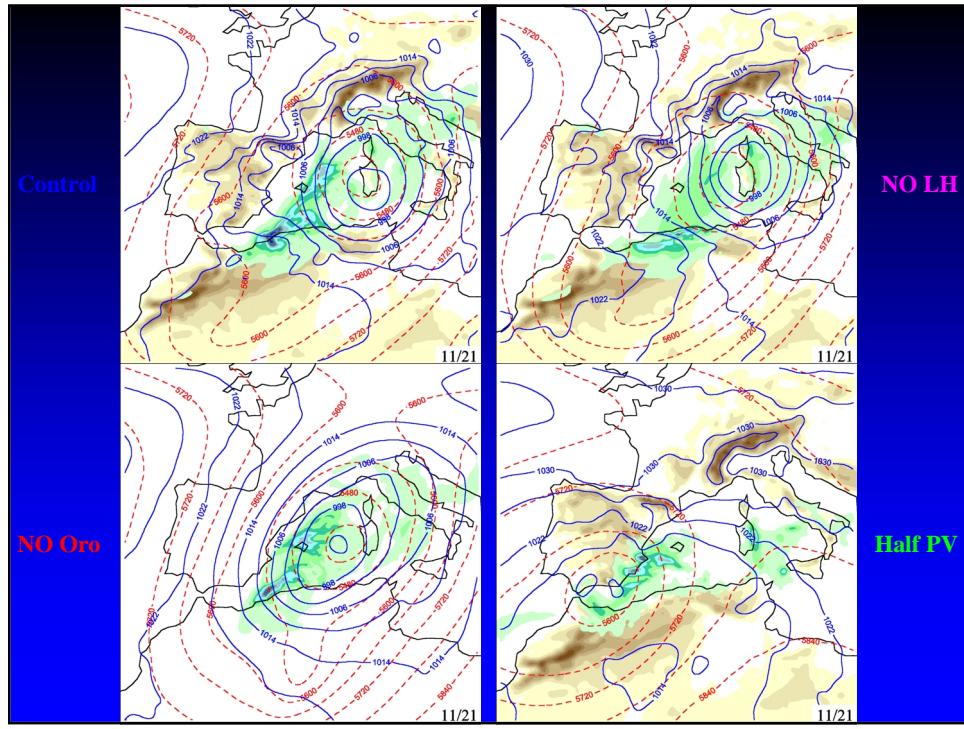
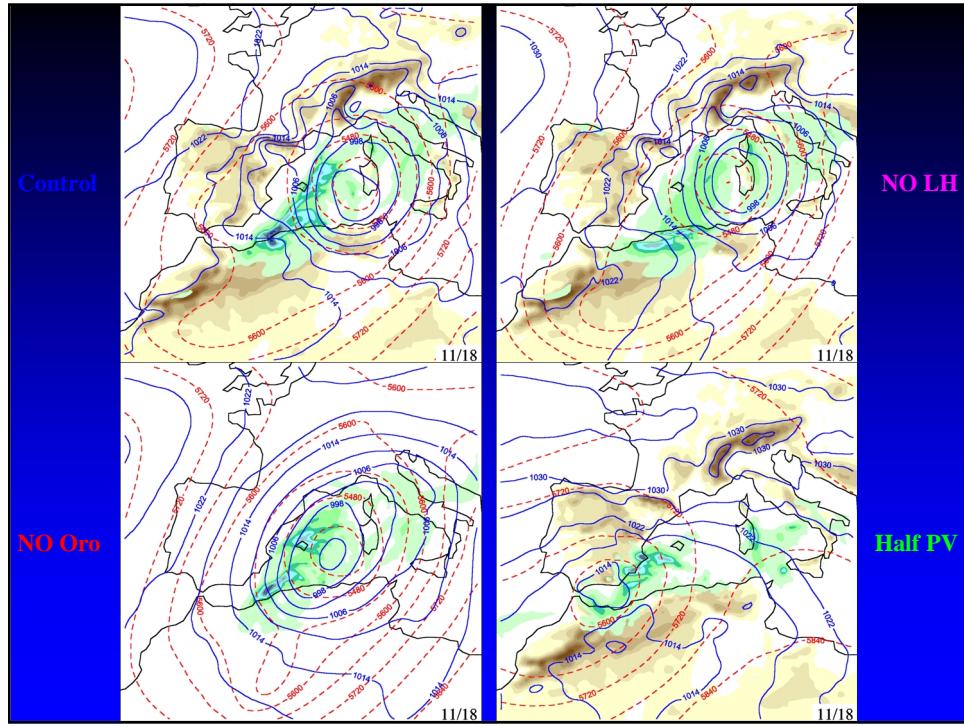


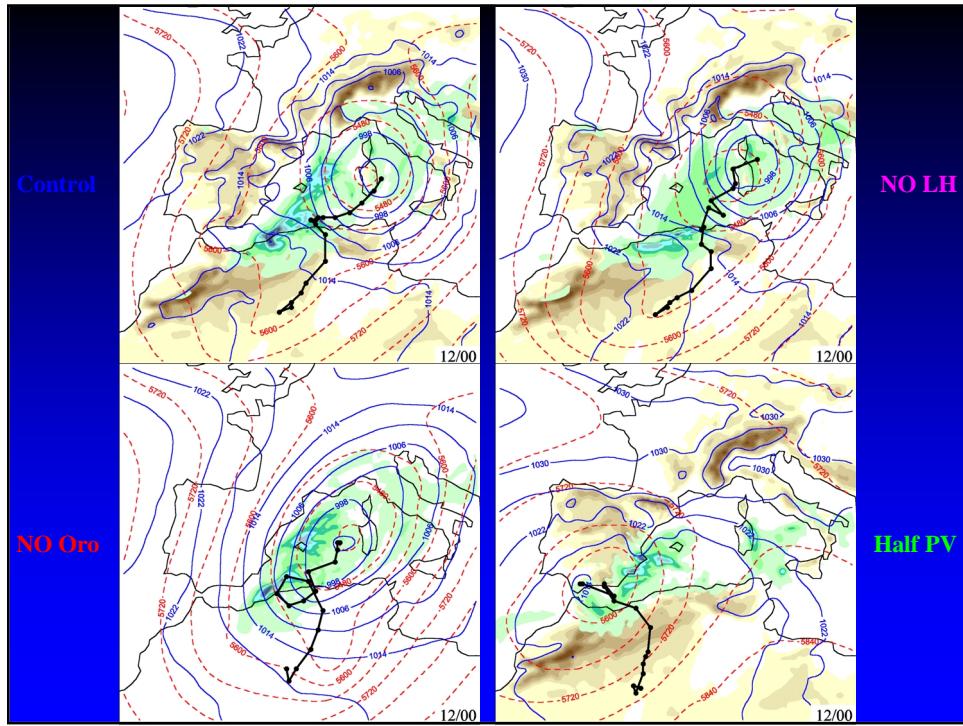
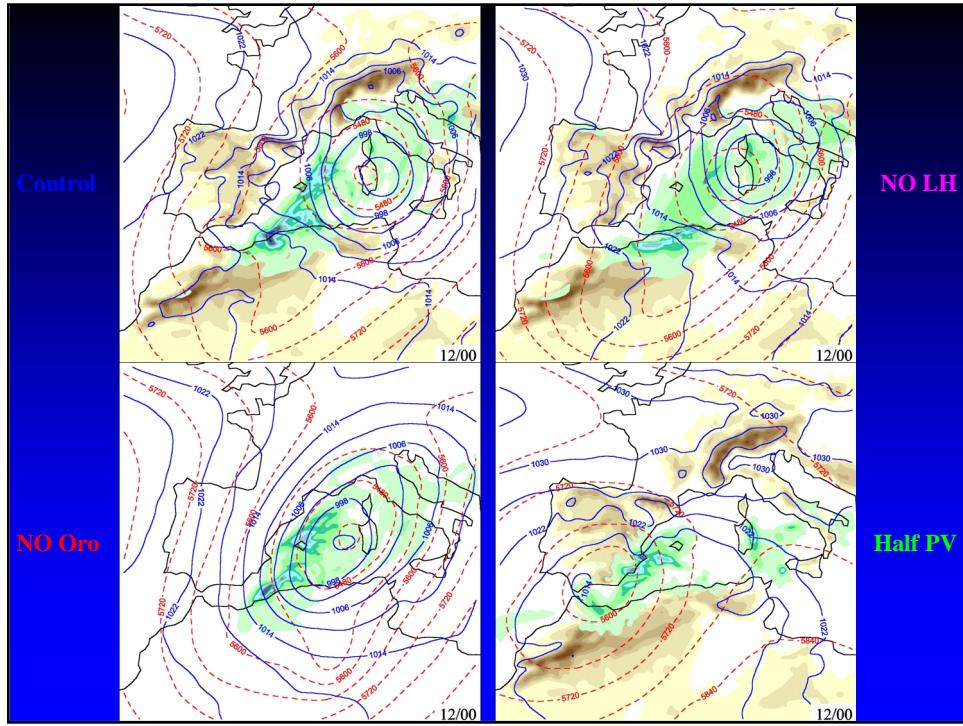


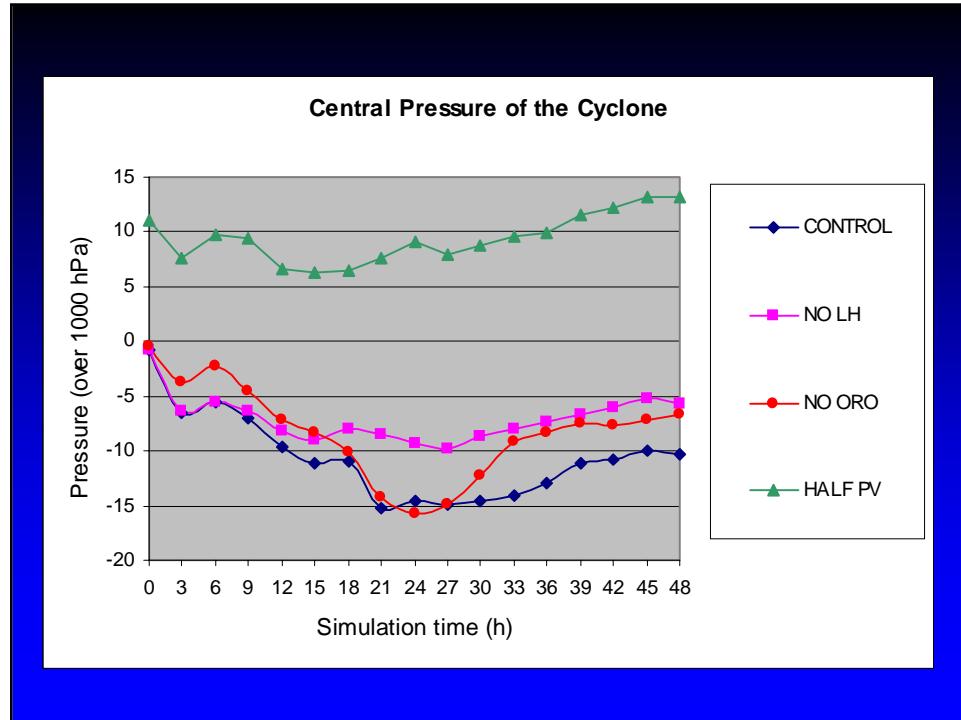












- ### SOME CONCLUSIONS
- ➡ Extreme cyclogenesis event in the western Mediterranean region (the worst storm affecting the Balearic Islands during the last decades)
 - ➡ Baroclinic development + Diabatic contribution from condensation
 - ➡ Quasi-geostrophic diagnosis:
 - Strong dynamical forcing for upward motion (and surface pressure fall)
 - Continuous moisture supply at low levels for heavy precipitation to occur
 - However, convective conditions farther eastward (within warm sector)
 - ➡ PV-based diagnosis:
 - Typical sequence of many extratropical cyclones: **Ulev → Llev → DIAB**
 - Controlled by the mutual interactions among the anomalies and mean flow
 - ➡ Numerical simulations:
 - Local orographic forcing was crucial for the flood-producing rainfall
 - **LH** and **Oro** modulated the deepening rate and trajectory of the cyclone